

Ecologia

ISSN 1996-4021



ISSN 1996-4021 / DOI: 10.3923/ecologia.2011.23.30

© 2011 Academic Journals Inc.

The White Cliffs of Dover are an Example of Natural Carbon Sequestration

Michael Witty

Department of Biology, Monmouth University, 400 Cedar Ave, West Long Branch, NJ 07764, USA

ABSTRACT

The purpose of this study is to review carbon sequestration by coccolithophores and consider it as a controllable technology. Recent research suggesting runaway global warming has inspired some scientists to propose methods to intervene in global geochemical processes and prevent future catastrophes. These include schemes to sequester carbon dioxide under oceans, use of orbiting mirrors to reflect energy into space before it can heat the atmosphere or seeding the atmosphere with sulfur compounds which precipitate clouds and reflect radiant energy. However, diverse scientific literature shows there is no agreement on the best methods to attempt and very little empirical experimentation to suggest which are practical. Additionally, there are some unexpected consequences suggested by critics which show many proposals for global geochemical engineering are fraught with danger and that they effectively replace one problem with another, rather than improve matters. However, geochemical systems are in stable equilibrium with each other and for that reason doing nothing may be the best tactic for global geochemical engineering. Geologic history has shown that natural processes give negative feedback to perturbations in the atmosphere, attenuating runaway global warming without human intervention and the White Cliffs of Dover are a great example of this. Control of calcite formation from coccoliths is a complex problem beyond our present technology.

Key words: Coccolithophore, coccolith, calcite, chalk, Dover

INTRODUCTION

During a scientific meeting at Monmouth University in April 2010 the subject of Global Warming (Ackerman and Knox, 2006) and its attenuation by carbon sequestration was considered. In particular the role of algae was discussed because it is strange but true that small creatures can have very large effects on ecosystems. One group is the coccolithophores (Green and Leadbeater, 1994) which are beyond clear observation by light microscopy but whose effects are very large because of their cell wall chemistry (Mukkamala et al., 2006). Coccolithophores are unicellular algae, common in oceanic ecosystems (Reid et al., 2009), which have a distinctive feature: cell walls composed of calcium carbonate plates (coccoliths, Fig. 1). Late in growth, blooms of these algae shed large amounts of their coccoliths. These turn water a milky turquoise color (Shutler et al., 2010) similar to the rock flour of glacial lakes (personal observation). This material is easily detected in satellite images because of its ability to backscatter visible light (Merico et al., 2003; Smyth et al., 2004) against a background of more transparent sea water. Their process of cell wall formation involves catalyzing the simple reaction of carbon dioxide (carbonic acid) to carbonate. This is a spontaneous reaction familiar to any scientist that has prepared a solution of sodium hydroxide then stored the solution in a bottle with a loose cap; a fine white precipitate of sodium carbonate

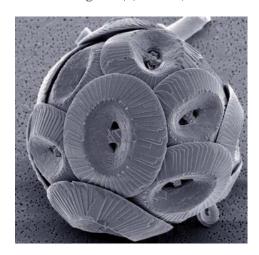


Fig. 1: Coccolithus pelagicus (Lampitt, 2008), showing button shaped cell wall plates composed of calcium carbonate. This species is common in cool North Atlantic waters (Cachao and Moita, 2000)

is formed by atmospheric CO_2 reacting with the NaOH. As with most life processes coccolith formation is an enzymic process and coccolithophores do not cause the reaction, but they do greatly increase the rate of reaction.

Much was learned about coccolith sedimentation and metamorphosis to calcite during the boom in North Sea Oil exploration of the 1960's and 1970's. Coccoliths are not significantly larger than the fine carbonate powders produced in chemical reactions and sediment at about 13.8 cm day⁻¹ (Steinmetz, 1994), so slowly that they may drift to unsaturated zones of the ocean which allow carbonate dissolution (Honjo, 1975) before fine carbonates can settle to the sea floor. However, because coccolithophore algae are consumed by animals such as copepods and aggregated into fecal pellets (Ziveria et al., 2007) they become sufficiently large to sediment rapidly (200 m day⁻¹, Steinmetz, 1994), accumulate and metamorphose to the mineral chalk, a form of cemented calcite particles (Neugebauer, 1974). This displaces the sea surface chemical equilibrium, drawing more CO_2 into reaction to carbonate. Rapid coccolith aggregate sedimentation is a process that has sequestered large amounts of CO_2 in the geological past (Gregory, 2002) and is a factor for enormous and rapid reduction of atmospheric CO_2 and therefore global warming (Fig. 2, Salter et al., 2007). Even so, coccoliths may suffer partial or complete dissolution after reaching ocean sediments (Roth and Berger, 1975) because deep sea water may not be saturated with carbonate.

These microscopic biological activities might be ignored by most people if it were not for their large scale effects that are visible to the unaided eye: for example, the White Cliffs of Dover (Fig. 2, Table 1). The cliffs are composed of fossilized coccolith aggregates (Saruwatari et al., 2008) that fell to the bottom of shallow seas and built up deposits which were mineralized to the chalky form of calcium carbonate (Mukkamala et al., 2006). This was an important natural process in warm shallow seas (Buitenhuis, 2000), when thick layers of calcite formed on European continental shelves when their margins were flooded (Katz et al., 2004; Kennedy and Garrison, 1975) during the cretaceous epoch (Tyson and Funnell, 1987). The White Cliffs are prominent, but they are just the tip of an iceburg of calcite deposits that extend over large areas of North Western Europe (Fig. 3) (Hakansson et al. (1991), Buitenhuis (2000) and citations within); naturally sequestered carbon, removed from the carbon dioxide rich atmosphere of the cretaceous epoch.



Fig. 2: The White Cliffs of Dover near the South Foreland Lighthouse, seen from St. Margaret's Bay, Dover, 51°8'6" N, 1°22'30" E (Jouan, 2006). The Lighthouse is 21 m high and the cliffs are 101 m high at that point (Clegg, personal communication)

Table 1: Physical features of carbon dioxide sequestered as calcite (De Foe and Compton 1925, Lowry 1930)

Parameter	Quantity	Unit	Note	Citation
Standard CaCO ₃ Density	2.71	$ m g~cm^{-3}$	Calcite	DeFoe and Compton (1925)
Density due to the CO_2 component of CaCO_3	2.13	${ m g~cm^{-3}}$		
Standard CO ₂ (g) Density	1.977	$\mathrm{g}\mathrm{L}^{-3}$	0°C	Lowry (1930)
	0.001977	${ m g~cm^{-3}}$	1 atm	
i.e., the CO_2 contained in calcite is 1077 times as	dense as CO ₂ (g))		
White Cliffs thickness (nominal)	100	m		Green and Leadbeater (1994)
Thickness of gas released over the White Cliffs	0	m	$0^{\circ}\mathrm{C}$	
If this amount of calcite reacted to $\mathrm{CO}_2\left(g\right)$	0	km	1 atm	
Troposphere thickness over Europe	17	km		Ackerman and Knox (2006)

GLOBAL ENGINEERING AND UNINTENDED CONSEQUENCES

Coccolithophores have long been recognized as unique and interesting algae (Siesser, 1994), but recently they have attracted interest from global engineering theorists. Ecologists with ambitions for global environmental engineering have looked at the White Cliffs of Dover and been inspired by the effects of coccolithophores (Woodward et al., 2009). Their ambition is to achieve similar large scale Carbon Sequestration and control atmosphere composition (Wilson and Gerard, 2007; De Baar et al., 2005). These scientists see the contribution by fossil fuels of 5.4×10¹² tonnes of carbon per year to the atmosphere as a problem that could be solved if the converse process by coccolithophores was induced i.e., consignment of 3.2×10¹² tonnes carbon per year as carbonate to sediments (Raven and Falkowski, 1999). These are comparable figures and might be used like tunable factors to balance one another and control atmospheric CO₂ content. The most commonly suggested method for achieving this process is stimulating blooms of coccolithophores by ocean fertilization with inorganic iron (De Baar et al., 2008; Jin et al., 2008) because iron is the limiting nutrient in oceanic ecosystems (Pomar and Hallock, 2008) i.e., creating Artificial White Cliffs of Dover by stimulating coccolithophore blooms. However several nutritional and other factors determine coccolithophore blooming (Tyrrell and Merico, 2004) and other algae

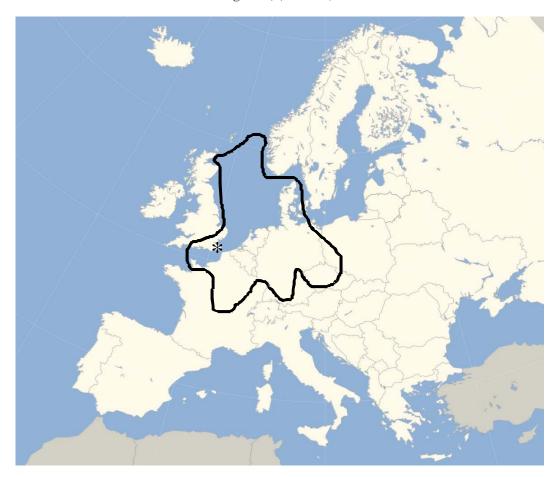


Fig. 3: Western Europe, showing the approximate extent of calcite deposits (Downing *et al.*, 1993) that formed as continental shelves were inundated (Tyson and Funnell, 1987; Hakansson *et al.*, 1991). The White Cliffs of Dover are indicated by "*". These deposits formed during the cretaceous epoch, a name derived from the Latin for chalk; *creta*

may also form blooms near the coccolith growth optima (Margalef, 1978; Balch, 2004), making the control of coccolithophore bloom processes tenuous. Observing field work is also limited by detecting coccolithophore bloom as scattered light: it is not only coccoliths that create significant back scatter (Tyrrell and Merico, 2004). Diatoms are significantly stimulated by iron fertilization (De Baar *et al.*, 2005) and have highly refractive silica cell walls. Therefore measurements that rely on satellites only must be treated with caution.

While the sight of the White Cliffs is inspiring it may be salutary to remember how common the unintended consequences of our actions are (Hu et al., 2003). Algal blooms, including blooms of coccolithophores are common, but how algae thrive or die and how they bloom is a generally unknown process and not easily manipulated by us in a controlled way (Marsh, 2003). Many of the methods of geochemical engineering have strong critics and increasing numbers of objections have been made (Robock, 2008; Robock et al., 2009; Hegerl and Solomon, 2009). This may be because of the very complex nature of both atmospheric and biological systems. For example, Raven and Falkowski (1999) list twenty two processes or factors involved in the movement of carbon from the atmosphere to oceans and sediments i.e., these are difficult systems to understand even if all

relevant factors are independent and linear (simple to extrapolate), or even known. Some particular unintended consequences of coccolithophore blooms are already known (Denman, 2008; Malin and Steinke, 2004) and include increased but unquantified denitrification and production of the long lived greenhouse gas N_2O (Law, 2008) i.e., plans to reduce carbon dioxide using coccolithophores would result in increases in other greenhouse gases. This is a paradox that has been seen before, for example carbon sequestration by increased forestation may, in some ecosystems, result in increased methane emission (Megonigal and Guenther, 2008), another greenhouse gas.

NATURAL PROCESSES ARE IN STABLE EQUILIBRIA

Part of the problem of global environmental engineering is an assumption that the climate is characterized by unstable equilibria which are capable of runaway changes to new stable states. However, if this were true the earth's climate could not recover from the many great knocks it has experienced in the past. Examples include volcanic eruptions which have covered millions of square kilometers in molten basalt (Pande, 2002) or darkened the atmosphere with large volumes of particulate matter (Schröder, 2002), which are later washed out by rain (Park et al., 2005). When the earth has been violently struck by comets there was an initial perturbation (Kring, 2007) then a resumption of our present desirable climate equilibrium, not a runaway and permanent disaster. In fact, asteroid impacts may be a regular occurrence in the history of the earth due to the passage of the solar system through galactic debris belts and life has always recovered (Bailer-Jones, 2009). These violent events indicate a stable rather than unstable equilibrium between factors controlling the climate; like a pendulum displaced from equilibrium, the atmosphere naturally swings back from extremes. The atmosphere is not like a tall vase which may teeter when knocked, then fall and occupy a second horizontal state of equilibrium, far from the first. If there's many a slip twixt cup and lip it is better to rely on natural processes, such as those that formed the White Cliffs of Dover and the stable equilibrium of geological processes we have all benefitted from for so long, than the uncertain outcomes of global engineering plans which have not been adequately formulated or tested.

ACKNOWLEDGMENTS

John Tiedemann and Dennis Rhoads inspired this study. Fang Xie of the Guggenheim library at Monmouth University was very helpful in preparation of this study.

REFERENCES

- Ackerman, S. and J.A. Knox, 2006. Meteorology. Brooks Cole Publishing Company, USA., ISBN-10: 0495108928.
- Bailer-Jones, C.A.L., 2009. The evidence for and against astronomical impacts on climate change and mass extinctions a review. Int. J. Astrobiol., 8: 213-219.
- Balch, W.M., 2004. Re-Evaluation of the Physiological Ecology of Coccolithophores. In: Coccolithophores, Thierstein, H.R. and J.R. Young (Eds.). Springer, New York, pp. 165-190.
- Buitenhuis, E.T., 2000. Interactions between *Emiliania huxleyi* and the Dissolved Inorganic Carbon System. University of Groningen, Netherlands.
- Cachao, M. and M.T. Moita, 2000. *Coccolithus pelagicus*, a productivity proxy related to moderate fronts off Western Iberia. Marine Micropaleontol., 39: 131-155.
- De Baar, H.J.W., L.J.A. Gerringa, P. Laan and K.R. Timmermans, 2008. Efficiency of carbon removal per added iron in ocean iron fertilization. Marine Ecol. Prog. Ser., 364: 269-282.

- De Baar, H.J.W., P.W. Boyd, K.H. Coale, M.R. Landry and A. Tsuda *et al.*, 2005. Synthesis of iron fertilization experiments: from the iron age in the age of enlightenment. J. Geophys. Res., 110: 24-24.
- De Foe, O.K. and A.H. Compton, 1925. The density of rock salt and calcite. Phys. Rev., 25: 618-620.
- Denman, K.L., 2008. Climate change, ocean processes and ocean iron fertilization. Marine Ecol. Prog. Ser., 364: 219-225.
- Downing, R.A., M. Price and G.P. Jones, 1993. The Making of an Aquifer: Chapter 1. In: The Hydrogeology of the Chalk of North-West Europe, Downing, R.A., M. Price and G.P. Jones (Eds.). Clarendon Press, Oxford, pp: 2.
- Green, J.C. and B.S.C. Leadbeater, 1994. The Haptophyte Algae. Oxford University Press, Oxford. Gregory, R.J., 2002. Carbon dioxide and climate over the past 300 Myr. Philosophical Trans. Royal Soc. Math. Phys. Eng. Sciences, 360: 659-673.
- Hakansson, E., R. Bromley and K. Perch-Nielsen, 1991. Maastrichtian Chalk of North-West Europe-a Pelagic Shelf Sediment. In: Pelagic Sediments on Land and Under the Sea, Hsu, K.J. and H.C. Jenkyns (Eds.). International Association of Sedimentologists Series, Wiley-Blackwell, New Jersey.
- Hegerl, G.C. and S. Solomon, 2009. Risks of climate engineering. Science, 325: 955-956.
- Honjo, S., 1975. Dissolution of Suspended Coccoliths in the Deep Sea Water Column and Sedimentation of Coccolith Ooze. In: Dissolution of Deep-Sea Carbonates, Sliter, W.V., A.W.H. Bé and W.H. Berger (Eds.). Cushman Foundation for Foraminiferal Research, Washington.
- Hu, J., N. Malic, J.L. Scott and C.R. Strauss, 2003. Potential impacts of deep-sea injection of CO₂ on marine organic chemistry. Green Chem., 5: 392-395.
- Jin, X., N. Gruber, H. Frenzel, S.C. Doney and J.C. McWilliams, 2008. The impact on atmospheric CO2 of iron fertilization induced changes in the ocean's biological pump. Biogeosciences, 5: 385-406.
- Jouan, R., 2006. South foreland lighthouse sur la falaise α Douvres. http://en.wikipedia.org/wiki/File:Douvres_(6).JPG.
- Katz, M.E., Z.V. Finkel, D. Grzebyk, A.H. Knoll and P.G. Falkowski, 2004. Evolutionary rajectories and biogeochemical impacts of marine eukaryotic phytoplankton. Annu. Rev. Ecol. Evol. Systematics, 35: 523-556.
- Kennedy, W.J. and R.E. Garrison, 1975. Morphology and genesis of hardgrounds and nodular chalks in the Upper Cretaceous of southern England. Sedimentology, 22: 311-386.
- Kring, D.A., 2007. The Chicxulub impact event and its environmental consequences at the Cretaceous–Tertiary boundary. Palaeogeography Palaeoclimatol. Palaeoecol., 255: 4-21.
- Lampitt, R., 2008. *Coccolithus pelagicus*. http://commons.wikimedia.org/ wiki/File: Coccolithus_pelagicus.jpg.
- Law, C.S., 2008. Predicting and monitoring the effects of large-scale ocean iron fertilization on marine trace gas emissions. Marine Ecol. Prog. Ser., 364: 283-288.
- Lowry, H.H., 1930. On the nature of active carbon. J. Phys. Chem., 34: 63-73.
- Malin, G. and M. Steinke, 2004. Dimethyl Sulfide Production: what is the Contribution of the Coccolithophores. In: Coccolithophores, Thierstein, H.R. and J.R. Young, (Eds.). Springer, Berlin, pp: 127-164.
- Margalef, R., 1978. Life-forms of phytoplankton as survival alternatives in an unstable environment. Oceanologica Acta., 1: 493-509.

- Marsh, M.E., 2003. Regulation of CaCO₃ formation in coccolithophores. Comp. Biochem. Physiol., 136: 743-754.
- Megonigal, J.P. and A.B. Guenther, 2008. Methane emissions from upland forest soils and vegetation. Tree Physiol., 28: 491-498.
- Merico, A., T. Tyrell, C.W. Brown, S.B. Groom and P.I. Miller, 2003. Analysis of satellite imagery for *Emiliania huxleyi* blooms in the Bering Sea before 1997. Geophys. Res. Lett., 30: 1337-1340.
- Mukkamala, S.B., C.E. Anson and A.K. Powell, 2006. Modelling calcium carbonate biomineralisation processes. J. Inorganic Biochem., 100: 1128-1138.
- Neugebauer, J., 1974. Some Aspects of Cementation in Chalk. In: Pelagic Sediments on Land and under the Sea, Hsu, K.J. and H.C. Jenkyns (Eds.). International Association of Sedimentologists, Blackwell Scientific Publications, Oxford, pp: 149-176.
- Pande, K., 2002. Age and duration of the deccan traps, India: A review of radiometric and paleomagnetic constraints. Proc. Indian Acad. Sci. Earth Planetary Sci., 111: 115-123.
- Park, R.J., D.J. Jacob, P.I. Palmer, A.D. Clarke and R.J. Weber *et al.*, 2005. Export efficiency of black carbon aerosol in continental outflow: Global implications. J. Geophys. Res., 110: 7-7.
- Pomar, L. and P. Hallock, 2008. Carbonate factories: A conundrum in sedimentary geology. Earth-Sci. Rev., 87: 134-169.
- Raven, J.A. and P.J. Falkowski, 1999. Oceanic sinks for atmospheric CO₂. Plant Cell Environ., 22: 741-755.
- Reid, P.C., A.C. Fischer, E. Lewis-Brown, M.P. Meredith and M. Sparrow *et al.*, 2009. Impacts of the oceans on climate change. Adv. Marine Biol., 56: 1-150.
- Robock, A., 2008. 20 Reasons why geoengineering may be a bad idea. Bull. Atomic Scientists, 64: 14-18.
- Robock, A., A. Marquardt, B. Kravitz and G. Stenchikov, 2009. The benefits, risks, and costs of stratospheric geoengineering. Geophys. Res. Lett., 36: 9-9.
- Roth, P.H. and W.H. Berger, 1975. Distribution and Dissolution of Coccoliths in the South and Central Pacific. In: Dissolution of Deep-Sea Carbonates, Sliter, W.V., A.W.H. Bé and W.H. Berger (Eds.). Cushman Foundation for Foraminiferal Research, Washington.
- Salter, I., R.S. Lampitt, R. Sanders, A. Poulton and A.E.S. Kemp *et al.*, 2007. Estimating carbon, silica and diatom export from a naturally fertilised phytoplankton bloom in the Southern Ocean using PELAGRA: A novel drifting sediment trap. Deep Sea Res., 54: 2233-2259.
- Saruwatari, K., N. Ozaki, H. Nagasawa and T. Kogure, 2008. Comparison of crystallographic orientations between living (*Emiliania huxleyi* and *Gephyrocapsa oceanica*) and fossil (*Watznaueria barnesiae*) coccoliths using electron microscopes. Am. Mineralogist, 93: 1670-1677.
- Schröder, W., 2002. The krakatoa event and associated phenomena. Earth Sciences History, 21: 199-206.
- Shutler, J.D., M.G. Grant, P.I. Miller, E. Rushton and K. Anderson, 2010. Coccolithophore bloom detection in the north east Atlantic using SeaWiFS: Algorithm description, application and sensitivity analysis. Remote Sensing Environ., 114: 1008-1016.
- Siesser, W.G., 1994. Historical Background of Coccolithophore Studies: Chapter 1. In: Coccolithophores, Winter, A. and W.G. Siesser (Eds.). Cambridge University Press, Cambridge.
- Smyth, T.J., T. Tyrrell and B. Tarrant, 2004. Time series of coccolithophore activity in the Barents Sea from twenty years of satellite imagery. Geophys. Res. Lett., 31: 4-4.

- Steinmetz, J.C., 1994. Sedimentation of Coccolithophores: Chapter 9. In: Coccolithophores, Winter, A. and W.G. Siesser (Eds.). Springer, UK.
- Tyrrell, T. and A. Merico, 2004. *Emiliania huxleyi*: Bloom Observation and the Conditions that Induce them. In: Coccolithophores, Thierstein, H.R. and J.R. Young (Eds.). Springer, New York.
- Tyson, R.V. and B.M. Funnell, 1987. European cretaceous shorelines, stage by stage. Palaeogeography Palaeoclimatol. Palaeoecol., 59: 69-91.
- Wilson, E. and D. Gerard, 2007. Carbon Capture and Sequestration Integrating Technology, Monitoring, Regulation. Wiley-Blackwell, New York.
- Woodward, F.I., R.D. Bardgett, J.A. Raven and A.M. Hetherington, 2009. Biological approaches to global environment change mitigation and remediation. Curr. Biol., 19: 615-623.
- Ziveria, P., B. de Bernardic, K.H. Baumannd, H.M. Stolle and P.G. Mortyna, 2007. Sinking of coccolith carbonate and potential contribution to organic carbon ballasting in the deep ocean. Deep Sea Res. Part II: Topical Stud. Oceanography 54: 659-675.