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Assessment of Compositional Heterogeneities of Core Samples from Harding (9/23b-10) and Gryphon (9/18b-17) Oil Columns

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Abstract: This study focuses on geochemical assessment of thirteen core samples from two wells (9/23b-10-Harding and 9/18b-17-Gryphon) that lie on the same block; so as to determine their compositional heterogeneities and to compare other biomarker parameters in these columns. Iatroscan TLC-FID showed that saturated, aromatic hydrocarbons, resins and asphaltene fractions were in agreement with where thermogenic hydrocarbons had existed. All the core samples were found to be associated with each other based on diagnostic biomarker parameters including Pr₁/Pr₁+Phy vs. C₂₇βα Dias/ C₂₇ sterane, Ts/Ts+Tm, 22S/22S+22R homohopane and the presence of 28,30 bisnorhopane from GC-MS analyses, which were related to their organic matter source and thermal maturity. This study reveals that vertical compositional heterogeneities remain in the two columns.

Key words: Compositional heterogeneities, biomarker, biodegradation, thermal maturity, core samples

INTRODUCTION

Geochemical analyses increase exploration efficiency by accounting for several parameters controlling the volumes of petroleum available for entrapment; this includes source rock quality and richness, thermal maturity and the timing of generation, migration and accumulation relative to trap formation; these require incorporation of other disciplines, such as seismic sequence stratigraphy and reservoir characterization (Peters and Fowler, 2002; Purvis *et al.*, 2002; Simoneit, 2004). Accurate and subsurface description is necessary to provide reservoir engineers with model for dynamic simulation (Purvis *et al.*, 2002; Wei *et al.*, 2008).

Large scale heterogeneities in petroleum composition within reservoirs have been known for many years but the topic received less attention until 1980s when sufficient analytical technologies became available to probe the phenomenon. The work of England *et al.* (1987) and England and Mackenzie (1989) has brought the explanation of petroleum reservoir geochemistry especially the concept of heterogeneity to the forefront of petroleum geochemistry (Stoddart *et al.*, 1995). These compositional differences are sometimes preserved on a lateral kilometre scale across communicating reservoirs over geological time but rarely within the vertical scale of a petroleum column (Stoddart *et al.*, 1995). The order of magnitude time scales for reservoir fluid mixing by diffusion as a function of reservoir properties was established by England and Mackenzie (1989). Homogenization times for inherited petroleum column

heterogeneities on a kilometre scale by lateral diffusion are on the order of 100 Ma while for vertical diffusion on a reservoir thickness scale of 100 m is about 1 Ma (England *et al.*, 1987; England and Cubit, 1995). For reservoirs filled during Late Tertiary times, some level of vertical homogenization would be expected in a petroleum column composition with persistence of a significant lateral heterogeneity (Larter and Aplin, 1995). Thus, these variations can be used to determine fill direction assuming that the most mature petroleum entering the reservoir last displaces the less mature petroleum (England *et al.*, 1987).

The variations in geochemical composition of petroleum fluids exist for one or more of the following reasons:

- The oils may be obtained from different source rocks, or may have differing contributions of oil from multiple source rocks as well as filling history (Killops and Killops, 1993; Tissot and Welte, 1984)
- Microbial alteration of petroleum during migration and emplacement (Baylis, 1998) and other processes including interaction of polar petroleum compounds with reservoir surfaces may affect homogenization even on small scales (Larter, 1992)
- Maturity differences may be preserved as compositional variations in a petroleum column due to incomplete mixing of petroleum in a reservoir (Horstad *et al.*, 1990)

Many development and production problems can be understood using rapid and common petroleum

geochemical methods (such as Rock eval pyrolysis, TLC-FID, GC, GC-MS, etc.). Traditional exploration focuses on subsurface traps and the play concept in sedimentary basins described according to tectonic style (Peters and Fowler, 2002; Wei *et al.*, 2008). Sophisticated analyses such as GC-IR-MS or Nitrogen compound and phenol analysis are found useful in the study of oils and core samples (Larter and Aplin, 1995).

This study is aimed at determining whether variations occur within and between the two oil columns located in Harding and Gryphon fields based on geochemical analyses associated with their maturity, biodegradation, source, etc. by analyzing oil saturated cores obtained from the exploration wells in the two fields.

MATERIALS AND METHODS

The samples used in this study were a gift from Dr. Martin of Newcastle University, upon Tyne, UK; comprising thirteen core samples: Eight samples from Harding oil column (9/23b-10) and five from Gryphon oil column (9/18b-17). The Gryphon oil field is located in block 9/18b, North-East of Aberdeen which was deposited during Late Palaeocene-Early Eocene; while the Harding field (Late Palaeocene-Early Eocene) lies principally across Block 9/23b North-East of Aberdeen. All reagents used in this study were of analytical grade except otherwise stated.

Iatroscan (Thin layer chromatography-flame ionization detector): Core extracts were analysed by Iatroscan TLC-FID. A TH-10 Mk IV instrument equipped with a flame ionization detector was used for rod scanning and quantification as described by Karlsen and Larter (1991). Silica rods type chromarods-S III (pore diameter 60', particle size 5 µm) were used for TLC. Normal hexane and toluene were used as mobile phases for the separation. The amounts of saturated and aromatic hydrocarbons and polar compounds (resins and asphaltenes) were identified by a digital recorder (LabSystem Atlas v3.0 software).

Thin Layer Chromatography (TLC): The TLC plates were prepared using aqueous slurry of 0.5 mm thick Keiselgel 60 G; the plates were activated for an hour at 100-120°C. The samples were spotted on separate TLC plates at about 2-3 cm from the bottom using a capillary tube. The plates were then separated in petroleum ether. Rhodamine 6 G was sprayed on the plates to identify different bands under UV light, aliphatic and aromatic hydrocarbon bands were identified and scraped off into different prepared elution tubes which were washed with about 40-50 mL

dichloromethane. The extract for each was reduced in volume using a rotary evaporator and then placed in auto-sample vial. Aliphatic hydrocarbon fractions were taken for further analysis.

Gas Chromatography-Mass Spectrophotometry (GC-MS):

GC-MS analyses of the aliphatic hydrocarbon fractions were performed on a Hewlett Packard 5890 gas chromatograph interfaced to a VG AutoSpecQ Ultima (electron energy 70 eV; electron multiplier 250 V; filament current 200 µA; source temperature 280°C) tuned to 1000 resolution. Chromatography was carried out on a fused silica column (30 m×0.25 mm i.d.) coated with DB5MS (modified 5% phenyl, 95% methyl silicone, 25 µm film thickness), using a splitless injection technique. The oven was programmed for an initial temperature of 40°C for 2 min, followed by heating at 4°C min⁻¹ to 310°C and held at a final temperature for 20 min. The samples were run using full scan and Single Ion Monitoring (SIM) and data obtained was recorded and processed using a HP ChemStation data system.

RESULTS AND DISCUSSION

Well 9/23b-10 (Fig. 1) showed approximately 15-90% saturated hydrocarbons, 10-49% aromatic compounds, 0.6-26% resins and 0-29% asphaltenes; also, well 9/18b-17 (Fig. 2) was relatively uniform with a slight increase in saturated hydrocarbons (e.g., from 45-68% at 5748-5782 m, respectively). The highest percentage of saturates in this well was observed at a depth of 5782 m, which has a very low concentration of resins and asphaltenes.

The plot of Pr₁/Pr₁+Phy versus C₂₇βα S+R Dias / C₂₇ steranes (Fig. 3) indicated a genetic relationship between the samples. Biomarker parameters such as 22S/22S+22R

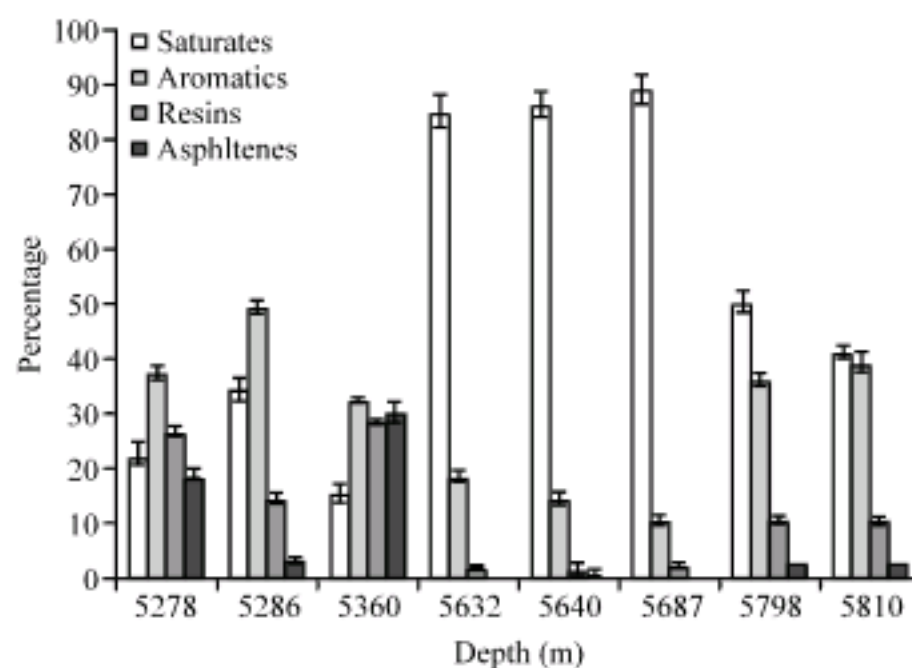


Fig. 1: Distribution of components of the core extracts in Harding oil column (9/23b-10) obtained from TLC-FID analyses

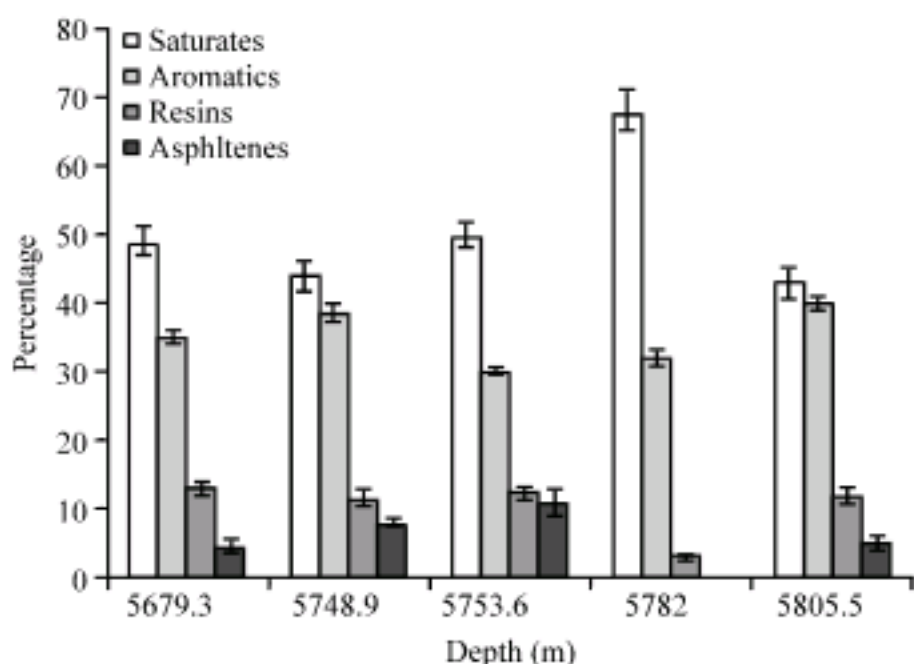


Fig. 2: Distribution of components of the core extracts in Gryphon oil column (9/18b-17) obtained from TLC-FID analyses

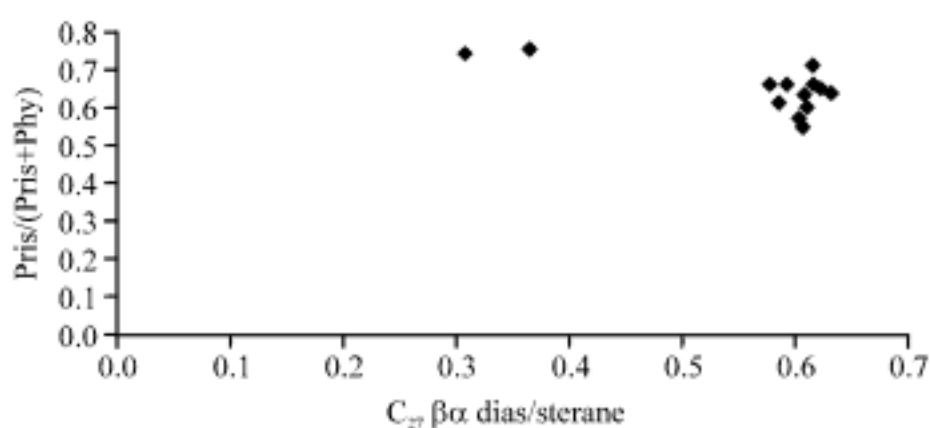


Fig. 3: Source related biomarker indicating the level of relationship between core samples in the two columns; this is obtained from GC-MS fingerprints

homohopane, Ts/Ts+Tm and sterane/hopane were measured for the two wells (Table 1) to determine the extent of these ratios in the two columns. Representative mass chromatogram (m/z 191) for hopane shows the presence of bisnorhopane (Fig. 4) in all the samples which is a source related parameter. Slight to moderate biodegradation of the samples occurred in both wells (9/23b-10 and 9/18b-17) as represented in Fig. 5a-d.

Utilizing the GC and GC/MS techniques made it possible to arrive at a clear characterization according to organic source materials, environmental conditions during deposition, the thermal maturity experienced by a rock or oil and the degree of biodegradation. This has been achieved using fingerprints from acyclic isoprenoids, steranes and terpanes biomarkers (El-Gayar *et al.*, 2002a; Bost *et al.*, 2001).

The gross compositions of core extracts screened by Iatrosan TLC-FID were determined; thus all samples in the two wells seem to represent regions where migrated petroleum has existed (Fig. 1, 2). This is due to the fact that the values obtained are higher than what were suggested by Bhullar *et al.* (1998) who showed that any

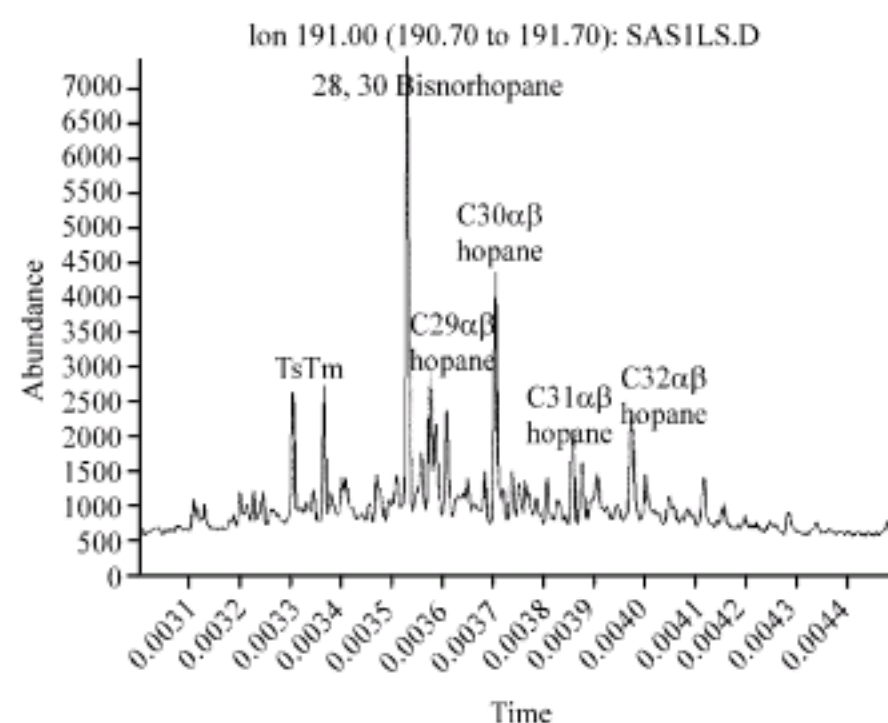


Fig. 4: Mass chromatogram (m/z 191) showing the presence of bisnorhopane

Table 1: Some of the diagnostic biomarker parameters in the two columns

Well depth (m)	Sample code	Ts/Ts+Tm	Sterane/hopane	Homohopane	
5278.00	9/23b-10	0.49	0.93	0.54	
5286.20		0.50	0.60	0.54	
5360.00		0.58	0.82	0.56	
5632.11		0.65	0.84	0.58	
5640.00		0.62	0.86	0.57	
5686.70		0.49	0.79	0.57	
5797.80		0.65	0.54	0.57	
5810.00		0.61	0.77	0.58	
5679.30		9/18b-17	0.69	0.49	0.59
5748.90			0.71	0.81	0.58
5753.60	0.69		0.50	0.59	
5782.00	0.72		0.61	0.60	
5805.60	0.69		0.50	0.59	

Where Ts/Ts+Tm ($C_{29} 17\alpha\text{-hopane}/C_{29} 17\alpha\text{-hopane} + C_{27} 17\alpha\text{-hopane}$); Sterane/hopane [$(C_{27}, C_{28}, C_{29} \text{ } \alpha\alpha\alpha (20S+20R) + C_{27}, C_{28}, C_{29} \text{ } \alpha\alpha\beta (20S+20R)) / (17\alpha\text{-hopane } C_{29}\text{-}C_{33} \text{ Pseudohomologs } 22S+22R \text{ for } C_{31}\text{-}C_{33})$] and homohopane [$(C_{35}/C_{31}\text{-}C_{35}) 17\alpha (22S+22R) \text{ hopane}$]. All of these were obtained from m/z 191 and m/z 217 chromatograms

core extract that contained less than 10% saturates, less than 20% aromatics and greater than 70% polar compounds had no thermogenic hydrocarbons and rather the extract was generated in-situ and does not represent a migrated petroleum. Some of the core samples especially those from 9/23b-10 showed the abundance of saturated hydrocarbons which are in agreement with suggested values (>60) of good saturated oils (Younes, 2003).

A genetic relationship between the core extracts is examined in Fig. 3; showing a convergence of points which indicates a common source with only two outliers; these outliers are the shallower samples of well 9/23b-10 which their fingerprints suggest that they were affected by biodegradation. This might likely be the reason that made them appeared different from the rest of the data because such difference is likely due to secondary processes rather than source.

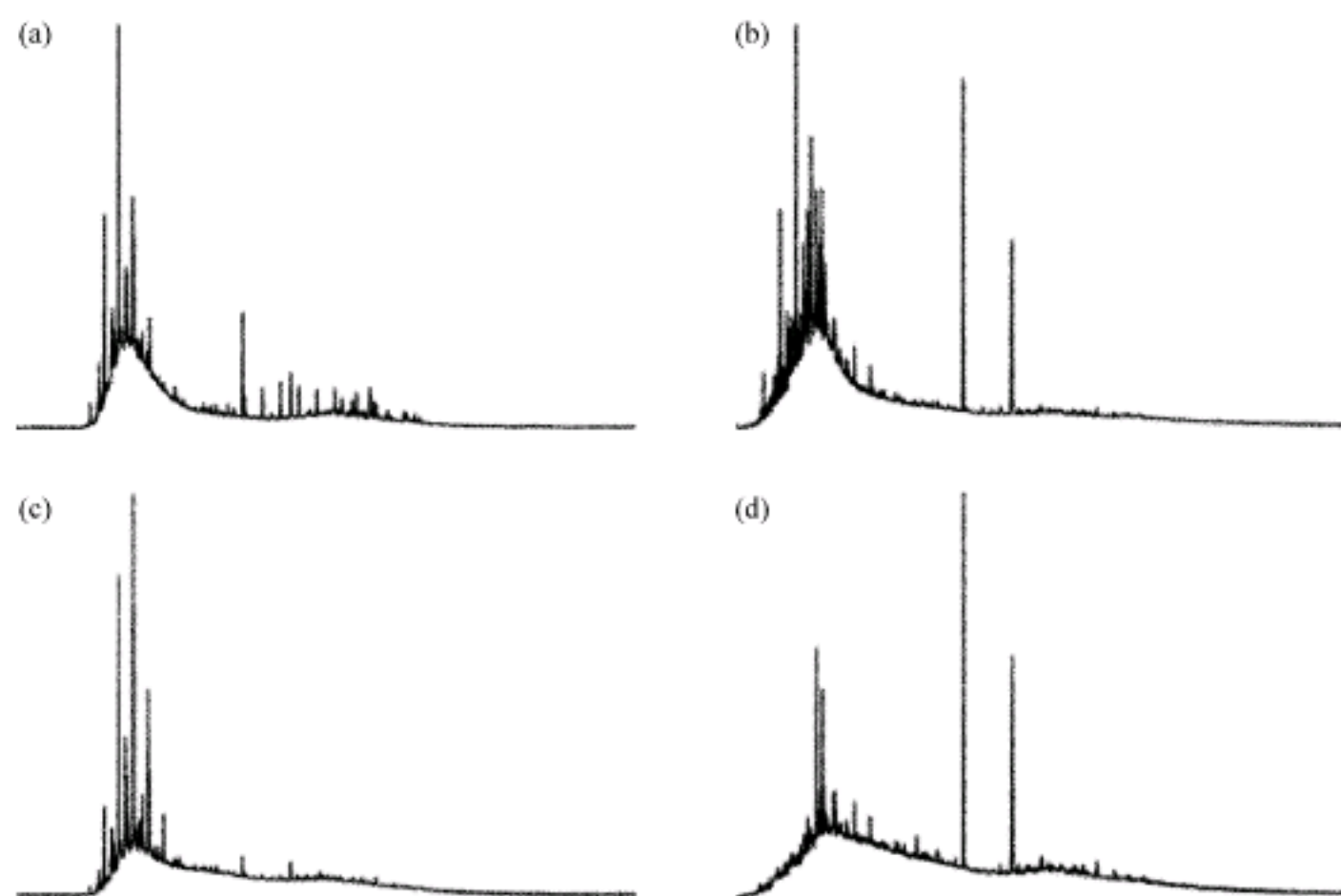


Fig. 5: Total hydrocarbon gas chromatograms indicating the presence of biodegraded oil and oil based mud signature from well-9/23b-10 (two at the top) and well-9/18b-17 (two at the bottom) from GC-MS analyses. (a) TIS: SAS2LSD (b) TIC: SAS15LSD (c) TIC: SAS12LSD and (d) TIC: SAS14LSD

The samples in both wells seem to be mature with higher values (Table 1), which conform to those described by Peters *et al.* (2005) where 22S/22S+22R homohopane ratios of 0.5-0.54 indicated a range of samples that have entered oil generation and those with 0.57-0.62 have reached the main phase of oil generation; they further added that once this stage has been reached, no further maturity information would be available because the ratio remains constant. This is in agreement with C₃₁ homohopanes measured for oil samples studied by El-Gayar *et al.* (2002b) and Harb *et al.* (2003); where all the values were closer or even higher than the steady-state value of 0.60 and as such the oils were suggested to have been derived from source rocks at advanced stage of thermal maturity.

Also, Waseda and Nishita (1998) described Ts/Ts+Tm ratio as one of the most reliable maturity indicator and the ratio increases from 0 to 1 with increasing maturity. Thus, the values obtained suggested some level of maturation in both wells. The Ts/Ts+Tm ratios obtained from this study were lower than those obtained from Western Desert samples that have higher values of ≥ 1 (Shi *et al.*, 2005; Harb *et al.*, 2003); and much higher than those from Paleozoic reservoir rocks in the Tataria and Perm basins, of Russia as found by Aizenshtat *et al.* (1998). Also, facies dependence could be exhibited by this ratio (Ts/Ts+Tm) where less than unity indicates marine depositional environment (Younes, 2003; Peters *et al.*, 2005).

Source related biomarker (sterane/hopane) has been described by Moldowan *et al.* (1985) where marine sources with major inputs from planktons and or benthic algae have sterane/hopane ratio ≥ 1 , while non marine sources with high terrestrial inputs tend to have very low values (near zero); therefore, all the values measured from this ratio are closer to 1, suggesting mixed inputs which may be linked to Kimmeridge source as the major source of oils in the North Sea (Dahl, 2004).

The occurrence of relatively high amounts of 28, 30-bisnorhopane in the m/z 191 mass chromatograms from most of the core samples in the two fields has been observed (Fig. 4). This feature seems to provide a good evidence for the source of these samples and tends to follow the findings from several studies which proposed the presence of bisnorhopane in the Viking Group, Kimmeridge Clay bitumen and oils from Oseberg Back Basin, Viking Graben, Heather Formation and Upper Jurassic Draupne Formation (Dahl, 2004). Thus 28, 30-Bisnorhopane found in the core extracts in this study indicates that the extracts were sourced from Kimmeridge Clay Formation. Thus, bisnorhopane isomers are found to be a helpful tool to group Viking Graben derived oils into families as represented in several petroleum systems and reservoir geochemistry analysis (Johannesen *et al.*, 2002).

Slight to moderate biodegradation occurred as most of the n-alkanes have been removed. These results agree with the study of Baylis (1998) and Peters and Fowler (2002) which showed that alteration of core materials and

contamination of samples with Oil-Based Mud (OBM) commonly give a saturated gas chromatogram with a hump of unresolved compounds; they further showed examples of hump in the C₁₃-C₁₆ n-alkane range where the sample was contaminated by Biovert (emulsion of water and highly paraffinic low aromatic mineral oil). This may be similar to what has been observed in Fig. 5. Also, the loss of straight-chain and isoprenoids as seen in the Fig. 5 corresponds to what was observed in Cuban oils studied by Grimalt *et al.* (2002); suggesting that mixing of severely biodegraded oil with undegraded oils during accumulation in the reservoir contributes to different patterns of biodegradation as well as oxidative activities of microbes within the reservoir (Campos *et al.*, 1996).

Based on these parameters; it could be concluded that all the core samples examined in this study, though obtained from the same block but are heterogeneous, with common source and compositional variations existed in the two columns indicating that vertical reservoir equilibration has not yet being established.

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