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## Simulation Study of past Climate Change Effect on Chickpea Phenology at Different Sowing Dates in Gorgan, Iran

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**Abstract:** This simulation study was mainly aimed to find the probable difference between sowing dates for response of four chickpea cultivars to past climate change. Firstly, the model CYRUS was recoded in QBASIC programming. Then phenology of cultivars Jam, Hashem, Arman and Beauvanij, seeded at day of year 50, 70 and 90, was evaluated during years 1961 to 2003 in Gorgan, Iran. The changes in some climatic variables were also studied. Results revealed that the solar radiation has been decreased for month December. Although the value of maximum temperature appeared to be the same across years 1961 to 2003 for all months, that of minimum temperature tended to show increasing trend for May and August. The increase in number of days with temperature higher than 35°C was considerable for April and October, but negligible for other months. The warming of night temperature was significant only for May and August. It found about 34 mm per 43 years decrease in monthly rainfall for March. The rate of increase in number of days with rainfall was 0.0737 day per year for December. Nearly similar situation was found for number of days with rainfall lower than and/or equal to 10 mm. The change in number of days with rainfall higher than 10 mm and lower and/or equal to 30 mm was in decreasing manner (March, October and December). The length of time from sowing to emergence appeared to be constant across past years. The advance in flowering ( $R_1$ ) was true only for cultivar Jam seeded at day of year 70. The length of period from  $R_1$  to pod initiation ( $R_3$ ) has been diminished for sowing at day of year 70 (Jam and Arman). Only cultivar Hashem seeded at day of year 90 tended to have decreasing trend for length of period between  $R_3$  and pod filling. It found one day (per 43 years) decreases from pod yellowing to maturity for cultivar Beauvamij seeded at day of year 70. Based on these findings and on the fact that future climate change is predicted using past changes, it seems that the adaptation strategies for future, including agronomy and/or breeding programs, may be not the same for different sowing dates of chickpea.

**Key words:** Phenology, chickpea, climate change, CYRUS, sowing date

### INTRODUCTION

There has been considerable concern in recent years about climatic changes caused by human activities and their effects on agriculture (Olfert and Weiss, 2006). Surface climate is always changing, but at the beginning of industrial revolution these changes have been more noticeable due to interference of human beings activity (Cutforth *et al.*, 1999). Studies of climate change impacts on agriculture initially focused on rising CO<sub>2</sub> levels (Curry *et al.*, 1990). Many researchers, including Pielke *et al.* (2002) reported that changes in temperature, radiation and precipitation need to be studied in order to evaluate the impact of climate change, beyond the CO<sub>2</sub> increase, on crops.

Although, there is general agreement for global warming, the rate of increase in temperature has been

different across the seasons and various regions. For example, the reports for Australia indicate that during the period between 1910 and 2000, average temperature has been increased by 0.76°C (McInnes *et al.*, 2002). In Iran, results regarding Kermanshah showed that except for January, February and March, the increase in temperature is considerable for all other months, especially June (Gholipoor *et al.*, 2006a). Lu *et al.* (2006b) reported that the increasing trend of temperature has been most considerable in the winter and early spring.

Temperature changes can affect crop productivity. Higher temperatures may increase plant carboxylation and stimulate higher photosynthesis, respiration and transpiration rates. Meanwhile, flowering may also be partially triggered by higher temperatures, while low temperatures may reduce energy use and increased sugar storage (Fiscus *et al.*, 1997). Reddy *et al.* (2002)

concluded that the rates of plant growth and development would continue to increase in the Southern US, because of enhanced metabolic rates at higher temperatures, combined with increased carbon availability. Changes in temperature can also affect air vapor pressure deficits, thus impacting the water use in agricultural landscapes (Kirschbaum, 2004). This coupling affects transpiration and can cause significant shifts in temperature and water loss. These feedbacks contribute to regional changes in precipitation and cloudiness, leading to changes in radiation (Pielke *et al.*, 2002).

The radiation has an important role in photosynthesis and crop productivity. There are more evidences for decreasing trend in solar radiation. For instance, Stanhill and Moreshet (1992) after analyzing data of 45 actinometrical stations for the years 1958, 1965, 1975 and 1985 indeed found a statistically significant worldwide decrease of global solar radiation averaging 5.3%. The decline was largest between 45 and 30°N. Regional declines have been also reported for the western as well as eastern sections of the former Soviet Union (Abakumova *et al.*, 1996).

It is globally accepted that precipitation is a leading factor affecting, especially, rainfed crops yield (Izaurrealde *et al.*, 2003). There are conflicting reports for change in precipitation during last century in different parts of the world. For example, in long-term mean precipitation, a decreasing trend of about  $-4.1 \text{ mm}^{-1} \text{ month}^{-1}$  100 years has been reported in boreal Asia. The largest and most statistically significant change has been a decline in rainfall in the winter-rainfall-dominated region of the far southwest of western Australia, where during the period of 1910-1995, winter (June-July-August) rainfall declined by 25%, mainly during the 1960s and 1970s (Smith *et al.*, 2000). In Central America, for much of the period from the early 1940s to 1995, Western Mexico has experienced an increasingly erratic monsoonal rainfall (Douglas and Englehart, 1999). Annual precipitation trends in past century are characterized essentially by enhanced precipitation in the northern half of Europe, with increases ranging from 10 to close 50% (Dore, 2005).

Generally, most studies are on fruit trees or trees and there are limited information on crop plant (Gholipoor *et al.*, 2006b). There is one published report regarding effect of past climate change on chickpea phenology (in Kermanshah, Iran) which has been based only on one sowing date (Gholipoor *et al.*, 2006b). This simulation study was mainly aimed to find the probable difference between sowing dates for response of four chickpea cultivars to past climate change; the changes in some climatic variables were also considered to be evaluated. The result of this study may be useful for selecting appropriate adapting strategy for sustaining the chickpea yield in future.

## MATERIALS AND METHODS

The weather data from 1961 to 2003 for Gorgan (36° 85' N, 54° 27' E and 100 m asl) Iran, was used for calculating the probable change in monthly values of some climatic variables. Data set contained daily values for sunshine hours, maximum temperature, minimum temperature and rainfall. Solar radiation data were calculated from sunshine hours and extraterrestrial solar radiation as outlined by Doorenbos and Pruitt (1977). A simple QBASIC program was written for calculating monthly values of some climatic variables including solar radiation, maximum and minimum temperatures, number of days with temperature greater than 35°C, number of days with temperature lower than 0°C, night temperature, day temperature, rainfall, number of days with rainfall, number of days with rainfall lower than and/or equal to 10 mm, number of days with rainfall greater than 10 and lower than and/or equal to 30 mm and number of days with rainfall greater than 30 mm. For calculating the rate of change in each climatic variable for period 1961-2003, the parameters a (intercept) and b (slope) in linear regression equation ( $Y = a + bX$ ) were calculated, using the Statistical Analysis System (SAS, 1989). The value of b was considered as rate of change.

We used the model CYRUS for investigating the phenological response of chickpea to past climate change in Gorgan, Iran. Firstly, it was recoded in QBASIC programming language and then run for above named weather data set. The cultivars were Hashem, Jam, Beauvanij and Arman and the sowing dates were 50, 70 and 90 day of year. In each year, the phenological phases including periods from sowing to emergence (sowing-E), from emergence to flowering (E-R<sub>1</sub>), from flowering to pod initiation (R<sub>1</sub>-R<sub>3</sub>), from pod initiation to pod filling (R<sub>3</sub>-R<sub>5</sub>), from pod filling to pod yellowing (R<sub>5</sub>-R<sub>7</sub>) and from pod yellowing to physiological maturity (R<sub>7</sub>-R<sub>8</sub>) were calculated. Like climatic variables, the value of b considered as rate of change in each phenological phase.

The model CYRUS was initially designed in 1999 by Soltani *et al.* (1999). Then it was developed for seedling emergence (Soltani *et al.*, 2006d), for leaf expansion and senescence (Soltani *et al.*, 2006c), for response of leaf expansion and transpiration to soil water deficit (Soltani *et al.*, 2000), for response to photoperiod (Soltani *et al.*, 2004a), for harvest index (Soltani *et al.*, 2005), for phenological development (Soltani *et al.*, 2006b) and for nitrogen accumulation and partitioning (Soltani *et al.*, 2006a). This model has been used for some simulation studies/investigations (Gholipoor and Soltani, 2005a, b, 2006; Gholipoor *et al.*, 2006a, b).

Briefly, in seedling emergence sub model of CYRUS, emergence response to temperature is described by a dent-like function with cardinal temperatures of 4.5 (base), 20.2 (lower optimum), 29.3 (upper optimum) and 40°C (ceiling temperature). Six physiological days (i.e., number of days under optimum temperature conditions) (equivalent to thermal time of 94°C days) are required from sowing to emergence at a sowing depth of 5 cm. The physiological days requirement is increased by 0.9 days for each centimeter increase in sowing depth. Snow cover effect is considered on the basis of daily maximum and minimum temperatures, as presented in Ritchie (1991).

In leaf sub model, cardinal temperatures for nod appearance are 6.0°C for base, 22.2°C for optimum and 31°C for ceiling temperature. Leaf senescence on the main stem starts when the main stem has about 12 nodes and proceeds at a rate of 1.67% per each day increase in physiological day (a day with non-limiting temperature and photoperiod). Leaf production per plant versus main stem node number occurs in two phases; phase 1 when plant leaf number increases with a slower and density-independent rate (three leaves per node) and phase 2 with a higher and density-dependent rate of leaf production (8-15 leaves per node).

Phenological development is calculated using multiplicative model that include a dent-like function for response to temperature and a quadratic function for response to photoperiod. Photoperiod-sensitivity is considered to be different in various cultivars and cardinal temperatures for phenological development are 0°C for base, 21°C for lower optimum, 32°C for upper optimum and 40°C for ceiling temperature. The cultivars require 25-31 physiological days from E to R<sub>1</sub>, 8-12 from R<sub>1</sub> to R<sub>3</sub>, 3-5 from R<sub>3</sub> to R<sub>5</sub>, 17-18 from R<sub>5</sub> to R<sub>7</sub> and 6 from R<sub>7</sub> to R<sub>8</sub> (Soltani *et al.*, 2006c).

The biomass production is calculated based on extinction coefficient (KS) and Radiation Use Efficiency (RUE). It assumes that KS is not radiation- and plant density-dependent. The RUE assumes to be constant (1 g MJ<sup>-1</sup>) across plant densities, but not across temperatures. After correction of RUE for temperature, it is not affected by either solar radiation or Vapor Pressure Deficit (VPD). The partitioning of biomass between leaves and stems is achieved in a biphasic pattern before first-seed stage. After this stage, the fixed partitioning coefficients are used for calculating biomass allocation.

Despite of many simulation models in which the linearity of harvest index increases has been used as a simple means to analyze and predict crop yield in experimental and simulation studies by Soltani *et al.* (2005) and related references for more detail), the CYRUS model assumes that its increase is biphasic with turning

point temperature equal to 17°C. The similar approach has been proved to be appropriate for application in wheat (Soltani *et al.*, 2004b).

The relation between total N and total biomass throughout the growth period is based on non-linear segmented model (with two segments/phases). Therefore, the rates of N accumulation during phase 1 and 2 are different and the turning point between two phases of N accumulation is considered 218.3 g biomass m<sup>-2</sup>. The distribution of N to different parts of plant is calculated using appropriate functions and coefficients.

In soil water balance sub model, daily soil water content is estimated as fraction transpirable soil water (FTSW, which ranges from 0 to 1) to calculate the degree of water limitation experienced by the crop. Similar to that described by Amir and Sinclair (1991), it accounted for additions from infiltration and losses from soil evaporation, transpiration and drainage. Infiltration is calculated from daily rainfall less any run-off. Run-off is estimated using the curve number technique (Knisel, 1980). Soil evaporation (Ev.) is calculated using the two-stage model as implemented in spring wheat model developed by Amir and Sinclair (1991). Stage I Ev. occurs when water present in the top 200 mm of soil and FTSW for the total profile is greater than 0.5. Stage II Ev. occurs when the water in the top layer is exhausted or the FTSW for the total soil profile reaches to less than 0.5. In stage II, Ev. is decreased substantially as a function of the square root of time since the start of stage II. The calculation of Ev. is returned to stage I only when rain or irrigation of greater than 10 mm occurs. Like procedure of Tanner and Sinclair (1983) and Sinclair (1994), the daily transpiration rate is calculated directly from the daily rate of biomass production, transpiration efficiency coefficient (= 5 Pa) and VPD. The calculation of VPD is based on suggestion of Tanner and Sinclair (1983) that it to be approximately 0.75 of the difference between saturated vapor pressure calculated from daily maximum and minimum temperatures.

## RESULTS AND DISCUSSION

The rate of change in solar radiation within the 43 years in Gorgan, Iran shown in Table 1. As it can be seen, the changes for solar radiation (MJ m<sup>-2</sup> day<sup>-1</sup> year<sup>-1</sup>) were negatively significant in December, but negligible in other months. The report of Stanhill and Moreshet (1992) for the years 1958 to 1985 indicates a statistically significant worldwide decrease of global solar radiation averaging 5.3% especially in North hemisphere. Similar results have been also reported for the western as well as eastern sections of the former Soviet Union

Table 1: The rate of change (value of b in equation  $Y = a+bX$ ) and its probability level (P) for solar radiation ( $\text{MJ m}^{-2} \text{ day}^{-1} \text{ year}^{-1}$ ), maximum temperature ( $T_{\text{max}}$ ;  $^{\circ}\text{C year}^{-1}$ ), minimum temperature ( $T_{\text{min}}$ ;  $^{\circ}\text{C year}^{-1}$ ) and number of days with temperature greater than  $35^{\circ}\text{C}$  ( $T>35$ ;  $\text{day year}^{-1}$ ) during 1961-2003

Month	Solar radiation		$T_{\text{max}}$		$T_{\text{min}}$		$T>35$	
	b	P	b	P	b	P	b	P
January	-0.0168	0.138	-0.0004	0.991	0.0024	0.917	---	---
February	-0.0155	0.198	0.0050	0.873	-0.0100	0.658	---	---
March	-0.0082	0.561	0.0028	0.922	-0.0125	0.502	---	---
April	-0.0233	0.270	0.0381	0.150	0.0079	0.672	<b>0.0115</b>	<b>0.011</b>
May	-0.0194	0.183	-0.0232	0.323	<b>0.0338</b>	<b>0.033</b>	-0.0281	0.106
June	0.0083	0.612	-0.0120	0.565	0.0012	0.926	-0.0424	0.255
July	-0.0261	0.178	-0.0127	0.494	0.0204	0.068	-0.0381	0.520
August	-0.0037	0.774	0.0017	0.923	<b>0.0345</b>	<b>0.009</b>	0.0252	0.666
September	-0.0041	0.727	0.0110	0.541	0.0188	0.204	0.0205	0.268
October	0.0094	0.396	0.0297	0.189	0.0106	0.594	<b>0.0165</b>	<b>0.003</b>
November	-0.0101	0.342	-0.0238	0.343	0.0043	0.823	-0.0005	0.812
December	<b>-0.0219</b>	<b>0.011</b>	-0.0248	0.301	-0.0167	0.484	---	---

The bold values are statistically significant

(Abakumova *et al.*, 1996). The natural and anthropogenic aerosols absorb solar radiation and this solar absorption within the atmosphere, together with the reflection of solar radiation to space, leads to a reduction in the solar radiation absorbed by the Earth's surface.

Although the changes in maximum temperature were not significant, but the trend was in increasing manner for February, March, April, August, September and October. Lu *et al.* (2006) and Fiscus *et al.* (1997) found the warming trend for the named temperature and demonstrated that, this can affect productivity of crops. The upwardly changes of minimum temperature appeared to be significant only for May ( $1.45^{\circ}\text{C}$  per 43 years) and August ( $1.48^{\circ}\text{C}$  per 43 years). Generally, reports indicating that globally averaged minimum temperatures continue to increase at a faster rate than the maximum temperatures, resulting in a narrowing of the diurnal temperature range (Karl *et al.*, 1993; Vose *et al.*, 2005). In Beijing, China it has been shown that the linear rate of increase in minimum temperature is  $4.08^{\circ}\text{C}/100$  year; whereas the maximum temperature increases with a linear rate of  $2.45^{\circ}\text{C}/100$  years (Karl *et al.*, 1993; Xie and Cao, 1996). Other reports for China indicating that in north regions, increasing trend of minimum temperature for period 1951-1990 has been more sensible, when compared to maximum temperature (Tao *et al.*, 2003). Analyzing daily and monthly maximum and minimum surface air temperatures at 66 weather stations over the eastern and central Tibetan Plateau for temporal trends also confirmed the asymmetric pattern of greater warming trends in minimum temperature or nighttime temperatures as compared to the daytime temperatures (Lu *et al.*, 2006). In central Europe, despite of low-lying stations for which the above named behavior has been found, data from mountain top stations show a similar increase for both minimum temperature and maximum temperature (Weber *et al.*, 1994). In Iran, results regarding Kermanshah showed that except for January,

February and March, the increase in minimum temperature is considerable for all other months, especially June (Gholipour *et al.*, 2006a); whereas maximum temperature appears to show increasing trend only for April, May, June and August; the increasing trend of temperature in Tabriz has tended to have asymmetric pattern too (Gholipour, 2008). Asymmetric change has been also reported for many other regions/countries, including southeastern Europe (Brázdil *et al.*, 1996), United States and Canada (Karl *et al.*, 1984). There are a number of possible factors, such as an increase in cloud cover, contributing to decreases in the diurnal temperature range (Easterling *et al.*, 1997; Lu *et al.*, 2006).

It found an increasing trend for number of days with temperature greater than  $35^{\circ}\text{C}$  in April and October, but statistically no change in other months (Table 1). The slope of this trend was more considerable for October as compared to April. The rate of change for number of days with temperature lower than  $0^{\circ}\text{C}$  and for night and day temperatures was shown in Table 2. Number of days with temperature less than  $0^{\circ}\text{C}$  tended to have no change across years 1961 to 2003. The rate of upwardly changes in night temperature ( $^{\circ}\text{C year}^{-1}$ ) was statistically considerable for May ( $1.37^{\circ}\text{C}$  per 43 years) and August ( $1.12^{\circ}\text{C}$  per 43 years), but negligible for others. The probability level for changes in day temperature tended to be higher than 0.05, suggesting no statistically changes in it.

The climate change has resulted in that the value of rainfall to be decreased only in March (Table 2); this is in agreement with Smith *et al.* (2000) who reported about  $4.1 \text{ mm year}^{-1}$  decrease in rainfall for years 1910 to 1995. The rate of increase in number of days with rainfall, say wet and/or rainy days, appeared to be  $0.0737 \text{ day year}^{-1}$  for December, but statistically zero for other months (Table 3); this increase in rainy days is contrary to report of Zhai *et al.* (1999). As shown in Table 3, there is similar

Table 2: The rate of change (value of b in equation  $Y = a+bX$ ) and its probability level (P) for number of days with temperature lower than 0°C ( $T < 0$ ; day year<sup>-1</sup>), night temperature (°C year<sup>-1</sup>), day temperature (°C year<sup>-1</sup>) and rainfall (mm year<sup>-1</sup>) during 1961-2003

Month	T<0		Night T		Day T		Rainfall	
	b	P	b	P	b	P	b	P
January	-0.0169	0.784	0.0016	0.948	0.0007	0.981	-0.3379	0.302
February	0.0083	0.866	-0.0063	0.790	0.0013	0.964	-0.1033	0.733
March	0.0257	0.211	-0.0081	0.683	-0.0011	0.964	<b>-0.7892</b>	<b>0.014</b>
April	---	---	0.0162	0.404	0.0309	0.191	-0.3563	0.250
May	---	---	<b>0.0319</b>	<b>0.050</b>	-0.0259	0.203	-0.2561	0.530
June	---	---	-0.0018	0.893	-0.0082	0.648	-0.2855	0.412
July	---	---	0.0120	0.312	-0.0042	0.790	-0.2619	0.320
August	---	---	<b>0.0260</b>	<b>0.042</b>	0.0098	0.520	-0.1022	0.733
September	---	---	0.0166	0.250	0.0135	0.403	0.1323	0.664
October	---	---	0.0157	0.425	0.0246	0.242	-0.7885	0.063
November	-0.0057	0.127	-0.0020	0.917	-0.0170	0.442	0.1945	0.694
December	0.0294	0.436	-0.0182	0.428	-0.0223	0.335	-0.5896	0.084

The bold values are statistically significant

Table 3: The rate of change (value of b in equation  $Y = a+bX$ ) and its probability level (P) for number of days with rainfall (wet days; day year<sup>-1</sup>), number of days with rainfall lower than and/or equal to 10 mm ( $R \leq 10$  mm; day year<sup>-1</sup>), number of days with rainfall greater than 10 and lower than and/or equal to 30 mm ( $10 < R \leq 30$  mm; day year<sup>-1</sup>) and number of days with rainfall greater than 30 mm ( $R > 30$  mm; day year<sup>-1</sup>) during 1961-2003

Month	Wet days		$R \leq 10$ mm		$10 < R \leq 30$ mm		$R > 30$ mm	
	b	P	b	P	b	P	b	P
January	0.0039	0.927	0.0169	0.665	-0.0151	0.337	0.0021	0.426
February	0.0104	0.785	0.0142	0.693	-0.0047	0.774	0.0009	0.734
March	-0.0071	0.863	0.0500	0.298	<b>-0.0571</b>	<b>0.003</b>	0.0000	1.000
April	-0.0142	0.752	-0.0088	0.839	-0.0076	0.551	0.0021	0.602
May	0.0403	0.304	0.0405	0.234	-0.0002	0.993	0.0000	1.000
June	0.0080	0.857	0.0044	0.918	0.0062	0.361	-0.0026	0.659
July	-0.0159	0.700	-0.0020	0.958	-0.0127	0.114	-0.0012	0.742
August	0.0051	0.886	0.0039	0.903	0.0027	0.784	-0.0015	0.731
September	-0.0148	0.690	-0.0308	0.387	0.0104	0.371	0.0056	0.253
October	-0.0026	0.948	0.0343	0.298	<b>-0.0272</b>	<b>0.050</b>	-0.0097	0.243
November	0.0455	0.251	0.0388	0.286	0.0023	0.886	0.0044	0.633
December	<b>0.0737</b>	<b>0.045</b>	<b>0.1108</b>	<b>0.005</b>	<b>-0.0314</b>	<b>0.041</b>	-0.0057	0.215

The bold values are statistically significant

Table 4: The rate of change (value of b in equation  $Y = a+bX$ ) and its probability level (P) for periods from sowing to emergence (sowing-E; day year<sup>-1</sup>), from emergence to flowering ( $E-R_1$ ; day year<sup>-1</sup>) and from flowering to pod initiation ( $R_1-R_3$ ; day year<sup>-1</sup>) for different cultivars and sowing dates during 1961-2003

Sowing date (day of year)	Cultivar	Sowing-E		$E-R_1$		$R_1-R_3$	
		b	P	b	P	b	P
50	Hashem	-0.0316	0.572	0.0787	0.392	-0.0003	0.951
	Jam	-0.0316	0.572	0.0735	0.415	0.0030	0.659
	Beauvanij	-0.0316	0.572	0.0711	0.435	0.0020	0.757
	Arman	-0.0316	0.572	0.0705	0.424	0.0032	0.551
70	Hashem	0.0008	0.972	-0.0575	0.363	-0.0042	0.385
	Jam	0.0008	0.972	<b>-0.0942</b>	<b>0.0318</b>	<b>-0.0191</b>	<b>0.050</b>
	Beauvanij	0.0008	0.972	-0.0584	0.357	-0.0015	0.809
	Arman	0.0008	0.972	-0.0569	0.365	<b>-0.0383</b>	<b>0.020</b>
90	Hashem	0.0038	0.785	0.0015	0.962	-0.0051	0.174
	Jam	0.0038	0.785	-0.0039	0.906	-0.0005	0.918
	Beauvanij	0.0038	0.785	0.0008	0.982	-0.0014	0.822
	Arman	0.0038	0.785	0.0035	0.910	-0.0009	0.853

The bold values are statistically significant

situation for number of days with rainfall lower than and/or equal to 10 mm ( $R \leq 10$  mm). Despite of these two variables, number of days with  $10 < R \leq 30$  appeared to show decreasing trend (March: 2.45 days per 43 years; October: 1.17; December: 1.35) and that of days with  $R > 30$  showed no change (Table 3).

Across the period 1961 to 2003, the time from sowing to emergence was statistically the same for all treatment combinations (i.e., sowing date × cultivar) (Table 4). The time from emergence to flowering ( $R_1$ ) has considerably been decreased for cultivar Jam sown at day of year 70 (4.0506 days per 43 years); the advancing of flowering is

Table 5: The rate of change (value of b in equation  $Y = a+bX$ ) and its probability level (P) for periods from pod initiation to pod filling ( $R_3$ - $R_5$ ; day year<sup>-1</sup>), from pod filling to pod yellowing ( $R_5$ - $R_7$ ; day year<sup>-1</sup>) and from pod yellowing to physiological maturity ( $R_7$ - $R_8$ ; day year<sup>-1</sup>) for different cultivars and sowing dates during 1961-2003

Sowing date (day of year)	Cultivar	$R_3$ - $R_5$		$R_5$ - $R_7$		$R_7$ - $R_8$	
		b	P	b	P	b	P
50	Hashem	-0.0045	0.395	-0.0074	0.372	0.0080	0.642
	Jam	-0.0023	0.702	0.0038	0.570	0.0005	0.985
	Beauvanij	0.0020	0.753	0.0002	0.984	0.0068	0.562
	Arman	-0.0029	0.478	0.0003	0.961	0.0063	0.641
70	Hashem	0.0011	0.794	0.0041	0.513	0.0118	0.582
	Jam	-0.0062	0.302	0.0082	0.214	-0.0039	0.845
	Beauvanij	0.0033	0.587	-0.0017	0.787	<b>-0.0234</b>	<b>0.0358</b>
	Arman	-0.0060	0.192	-0.0086	0.159	0.0051	0.775
90	Hashem	<b>-0.0377</b>	<b>0.040</b>	-0.0023	0.717	-0.0130	0.515
	Jam	0.0085	0.171	-0.0009	0.895	-0.0124	0.466
	Beauvanij	0.0030	0.616	0.0018	0.784	-0.0069	0.647
	Arman	-0.0038	0.369	0.0030	0.628	-0.0159	0.359

The bold values are statistically significant

in agreement with reports of the Menzel and Fabian (1999) and Lu *et al.* (2006) for other species; the change in length of this period was statistically negligible for other treatment combinations. The length of period between flowering and pod initiation ( $R_3$ ) appeared to show decreasing trend only for sowing at day of year 70 (cultivar Jam: 0.8213 day par 43 years; cultivar Arman: 1.6469 day par 43 years). The decreasing rate of time from pod initiation to pod filling ( $R_5$ ) was 0.0377 day per year (1.6211 day per 43 years) for cultivar Hashem sown at day of year 90, but statistically zero for other cultivars sown at this date and at other dates (Table 5). The length of period between pod filling and pod yellowing ( $R_7$ ) appeared to be the same across 43 years for all sowing dates and cultivars (Table 5). The decreasing trend regarding length of stage pod yellowing to stage physiological maturity ( $R_8$ ) appeared to be considerable for cultivar Beauvanij which seeded at day of year 70, but negligible for other treatment combinations. The climate-change-related advance in some phenophases of chickpea is in agreement with reports for other plants/crops (Fitter *et al.*, 1995; Sparks *et al.*, 2000; Lu *et al.*, 2006). As it was speculated by Menzel and Fabian (1999) the decrease in length of phenophases and consequently increasing the rest of growing season, may improve the opportunity for sowing the succession crop(s).

### CONCLUSION

Generally, the result of this study revealed that the solar radiation has been decreased for month December. Although the value of maximum temperature appeared to be the same across years 1961 to 2003 for all months, that of minimum temperature tended to show increasing trend for May and August. The increase in number of days with temperature higher than 35°C was considerable for April and October, but negligible for other months. The

warming of night temperature was significant only for May and August. It found about 34 mm per 43 years decrease in monthly rainfall for March. The rate of increase in number of days with rainfall was 0.0737 day year<sup>-1</sup> for December. Nearly similar situation was found for number of days with rainfall lower than and/or equal to 10 mm. The change in number of days with rainfall higher than 10 mm and lower than and/or equal to 30 mm was in decreasing manner (March, October and December).

The length of time from sowing to emergence appeared to be constant across past years. The advance in flowering ( $R_1$ ) was true only for cultivar Jam seeded at day of year 70. The length of period from  $R_1$  to pod initiation ( $R_3$ ) has been diminished for sowing at day of year 70 (Jam and Arman). Only cultivar Hashem seeded at day of year 90 tended to have decreasing trend for length of period between  $R_3$  and pod filling. It found one day per 43 years decrease from pod yellowing to maturity for cultivar Beauvanij seeded at day of year 70.

Because of highly difference between months for climate change, the difference between sowing dates for climate-related hastening in chickpea phenophases appeared to be considerable. Based on this finding and on the fact that future climate change is predicted using past changes, it seems that the adaptation strategies for future, including agronomy and/or breeding programs, may be not the same for different sowing dates of chickpea.

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